

How solar heating of the lower stratosphere influences tropospheric climate

Joanna D. Haigh

Space and Atmospheric Physics, Imperial College London, UK

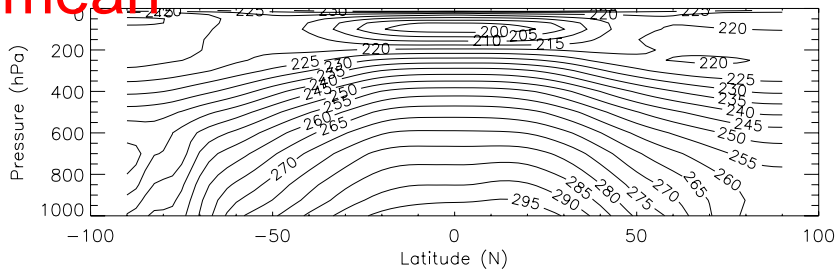
Michael Blackburn

Department of Meteorology, University of Reading, UK

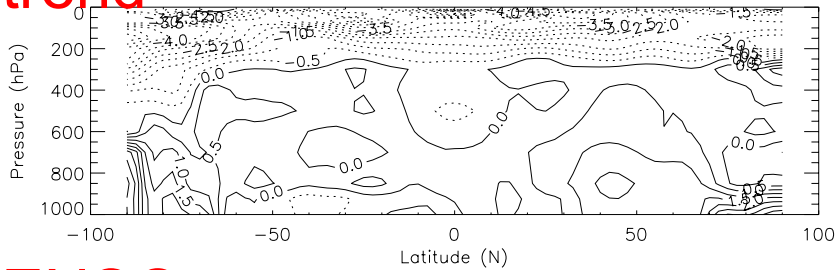
SORCE meeting Meredith NH, USA 27-29 October 2004

Multiple regression analysis: NCEP/NCAR zonal mean temperature

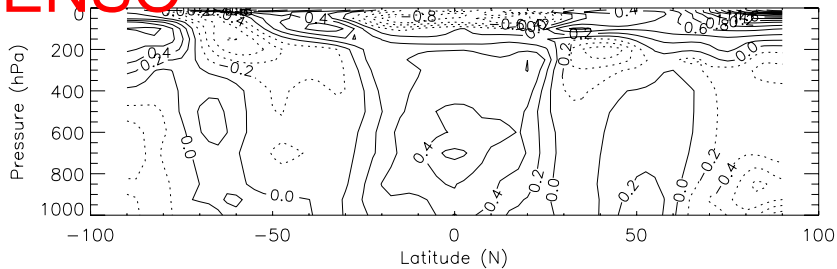
mean



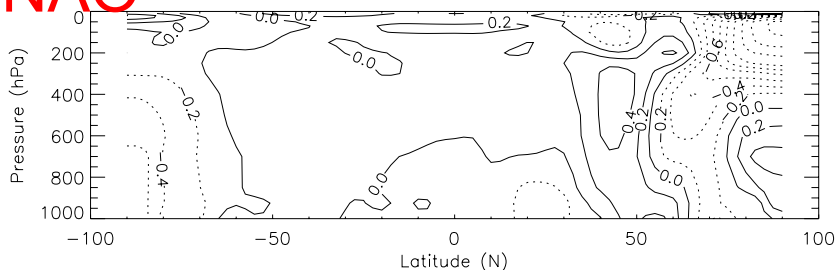
trend



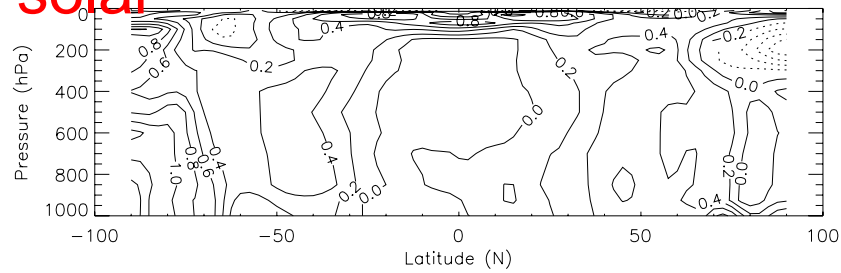
ENSO



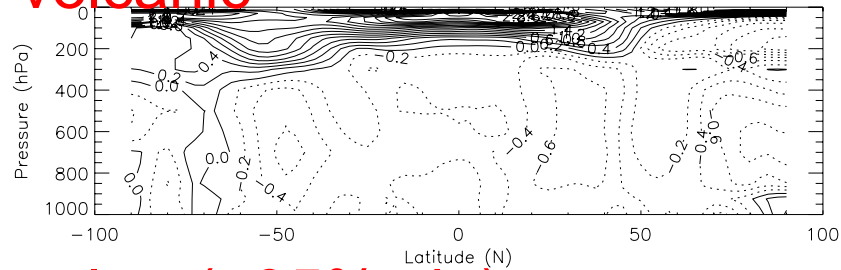
NAO



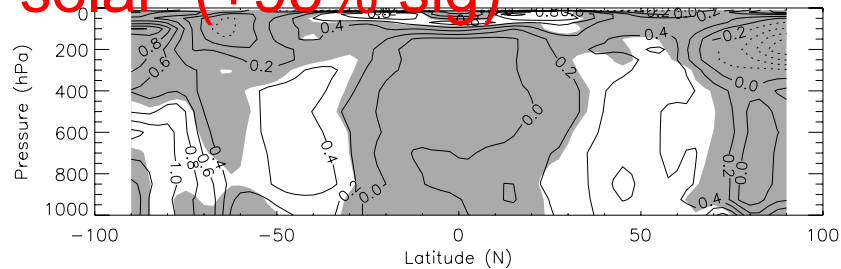
solar



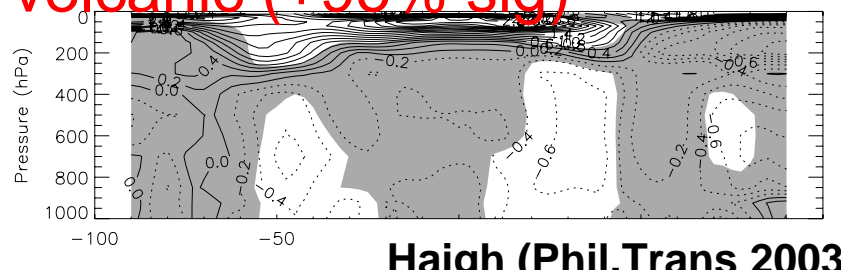
volcanic



solar (+95% sig)



volcanic (+95% sig)



Solar influence on zonal mean zonal wind

GCM simulation.

(a) Control run

(b) Difference between solar maximum and solar minimum.

Haigh (*Science*, 1996)

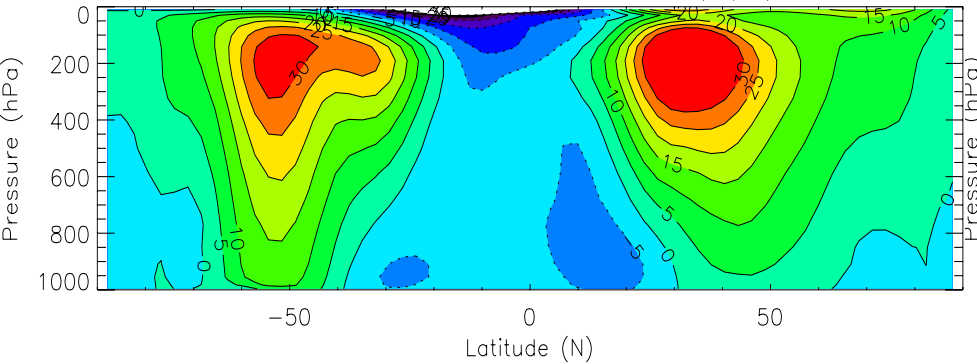
Multiple regression analysis of NCEP Reanalysis data 1979-2002.

(a) Mean

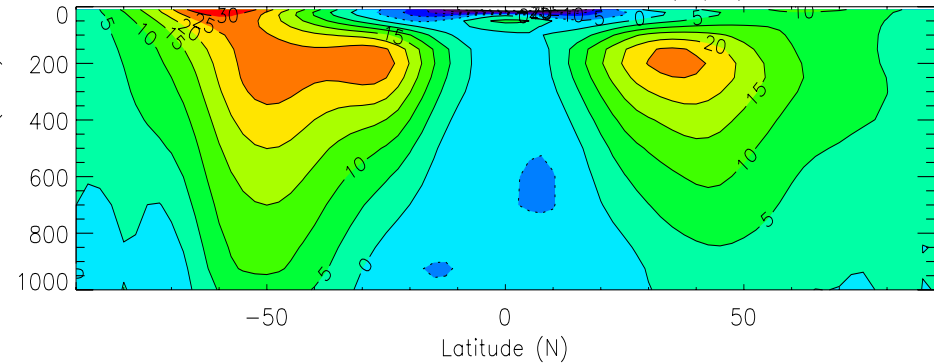
(b) Component associated with solar 11-year cycle.

Haigh et al (*J.Clim.*, submitted 2004)

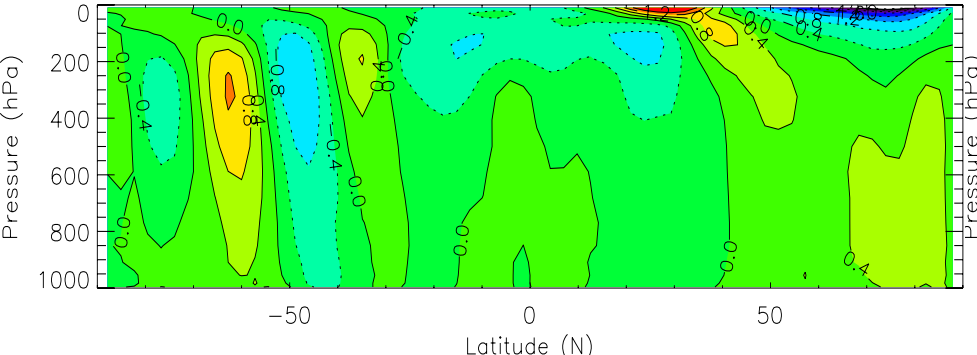
UGCM zonal mean zonal wind (m/s)



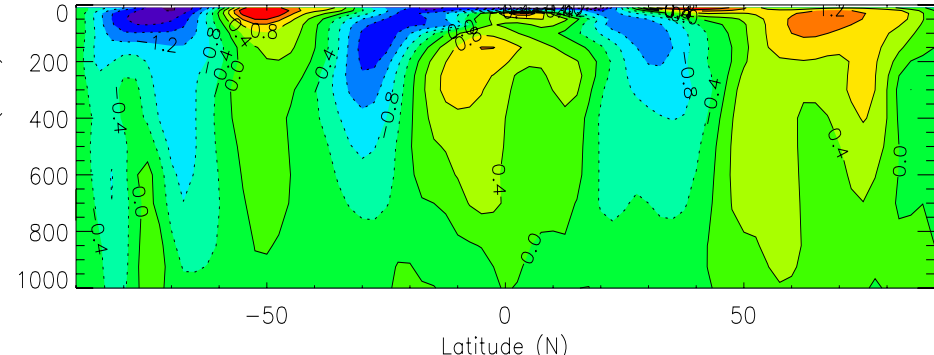
NCEP zonal mean zonal wind (m/s)



UGCM zonal mean zonal wind solar max - solar min (m/s)



NCEP zonal mean zonal wind solar max - solar min (m/s)



Potential mechanisms

- Increased static stability weakens tropical upwelling and Hadley cells
- Changes in u affect growth/propagation of planetary waves
- Baroclinic lifecycles modified

Simplified GCM experiments

Dynamical core of University of Reading spectral model:

Full dynamics T42 L20. No orography.

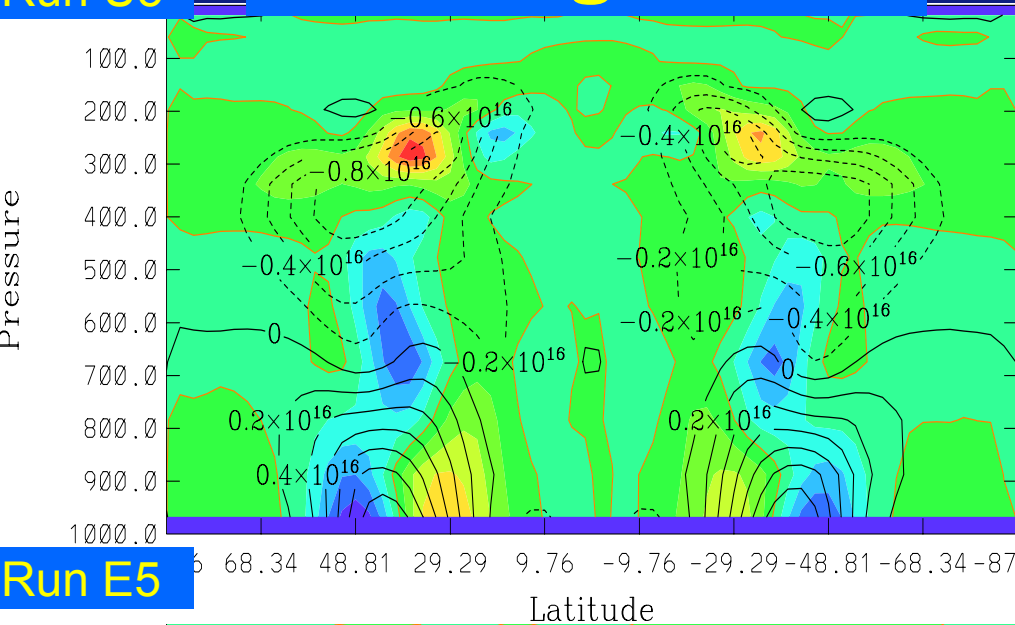
Newtonian cooling (equinoctial radiative equilibrium temperatures).

Rayleigh friction.

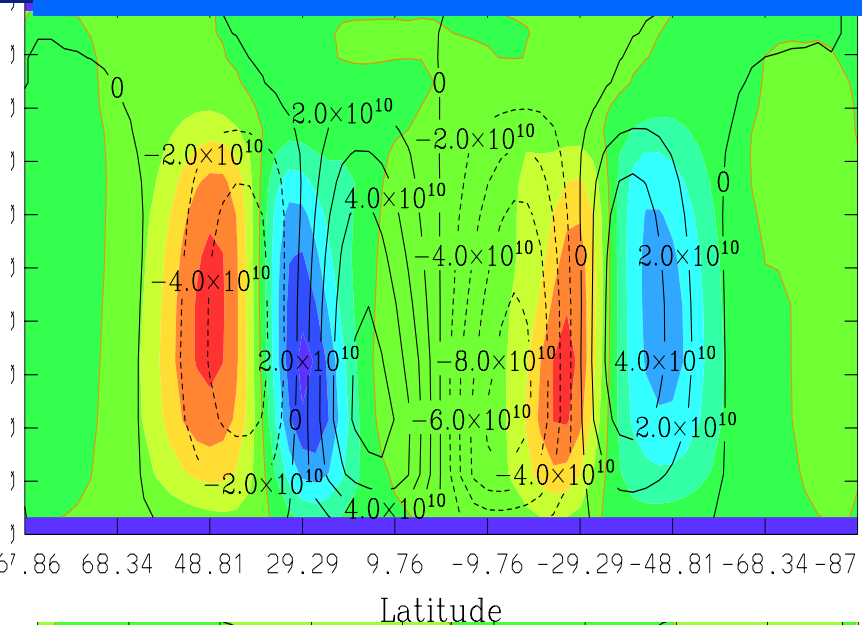
Run C	Control: T_e distribution of Held and Suarez (1994).	10,000 days
Run U1	Stratospheric T_e (ONLY) increased uniformly by 1K.	10,000 days
Run E1	Stratospheric T_e (ONLY) increased by 1K at the equator, decreasing with $\cos^2(\text{latitude})$ to 0K at the poles.	10,000 days
Run U5	5K uniform increase in stratospheric T_e	1,000 days
Run E5	5K $\cos^2(\text{lat})$ increase in stratospheric T_e	1,000 days

E-P flux divergence

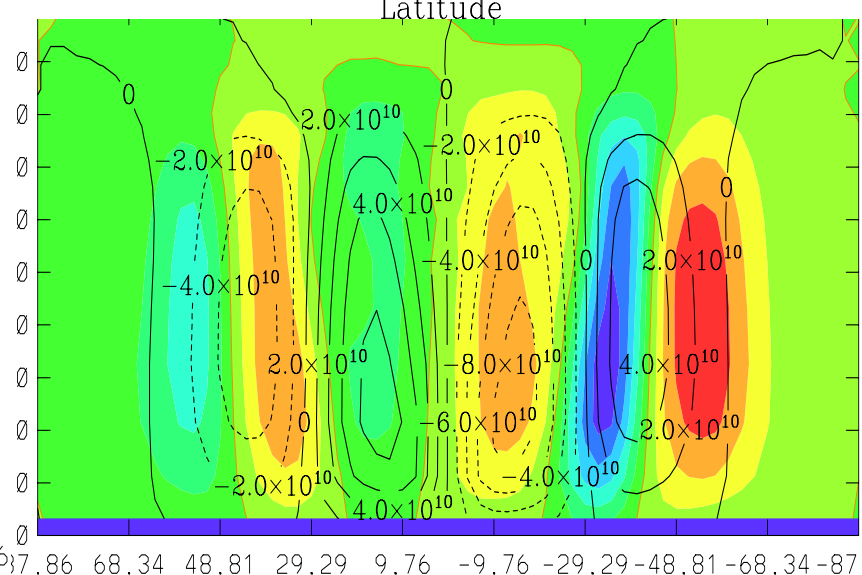
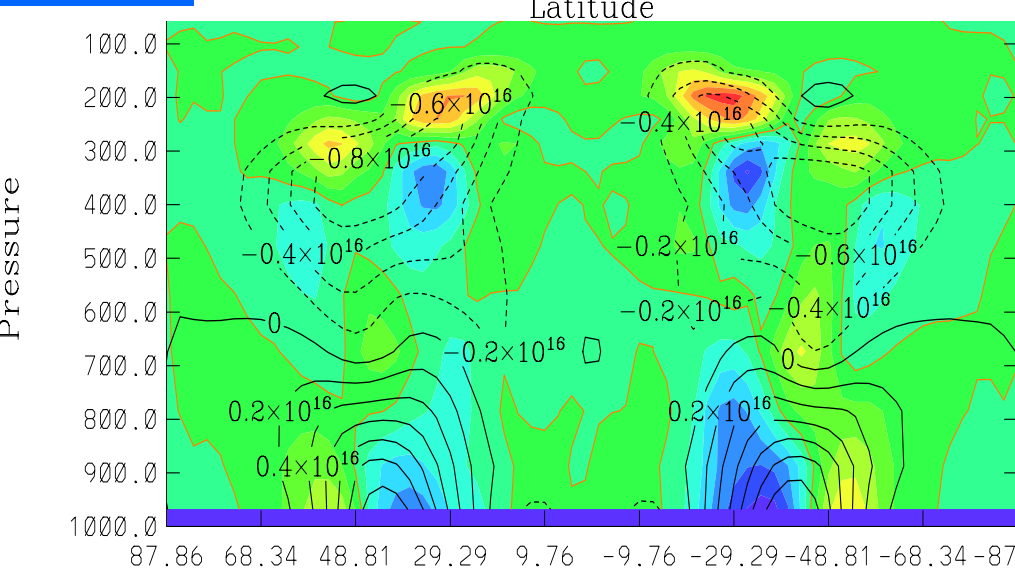
Run U5



Mean meridional circulation



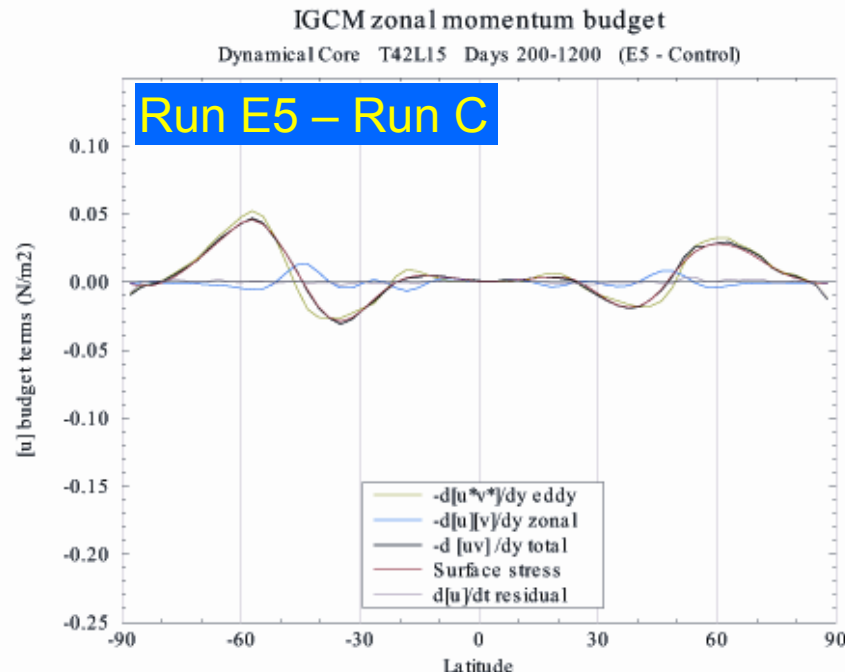
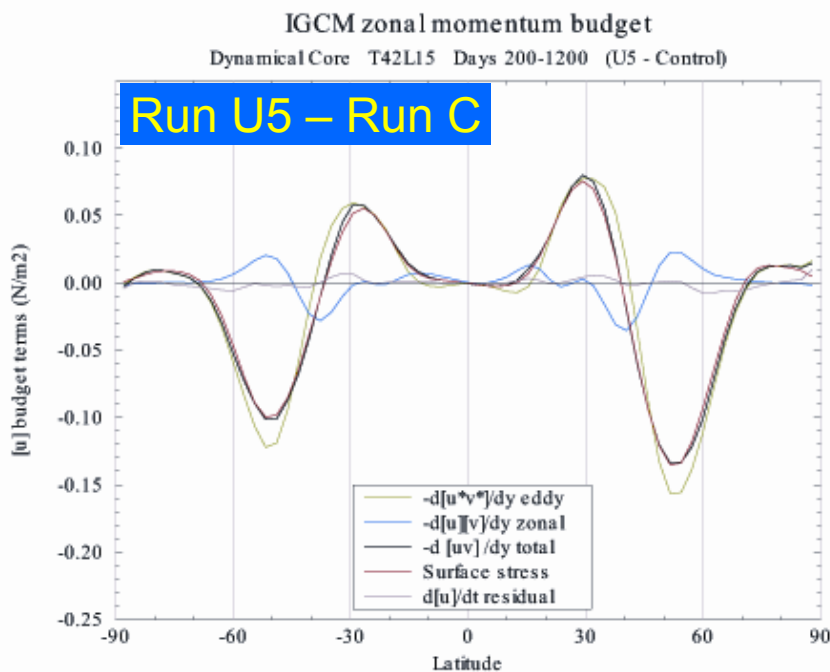
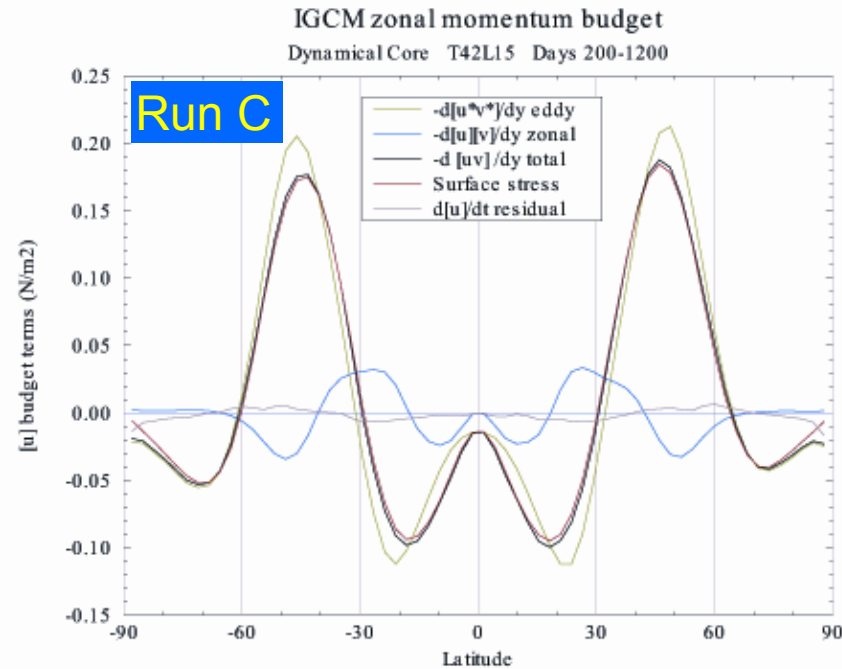
Run E5



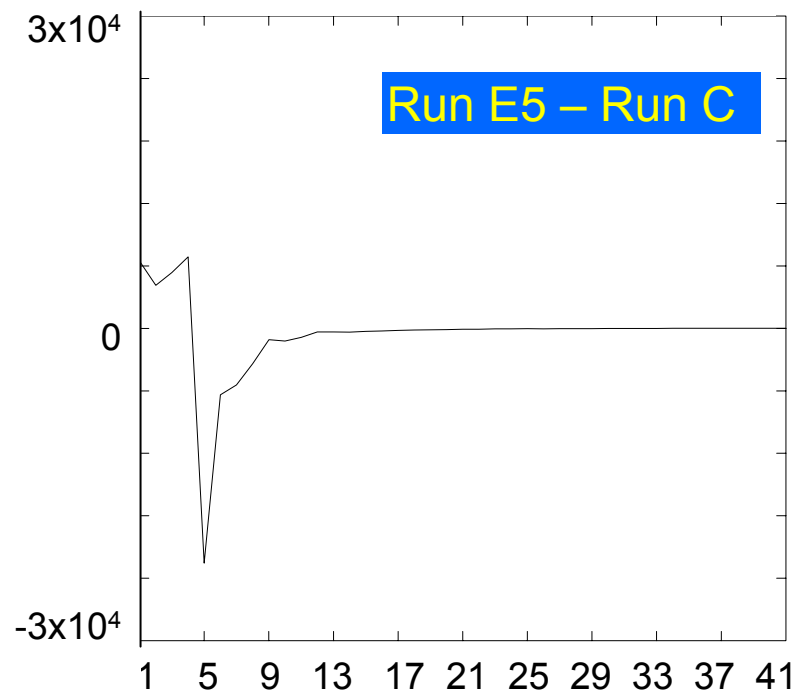
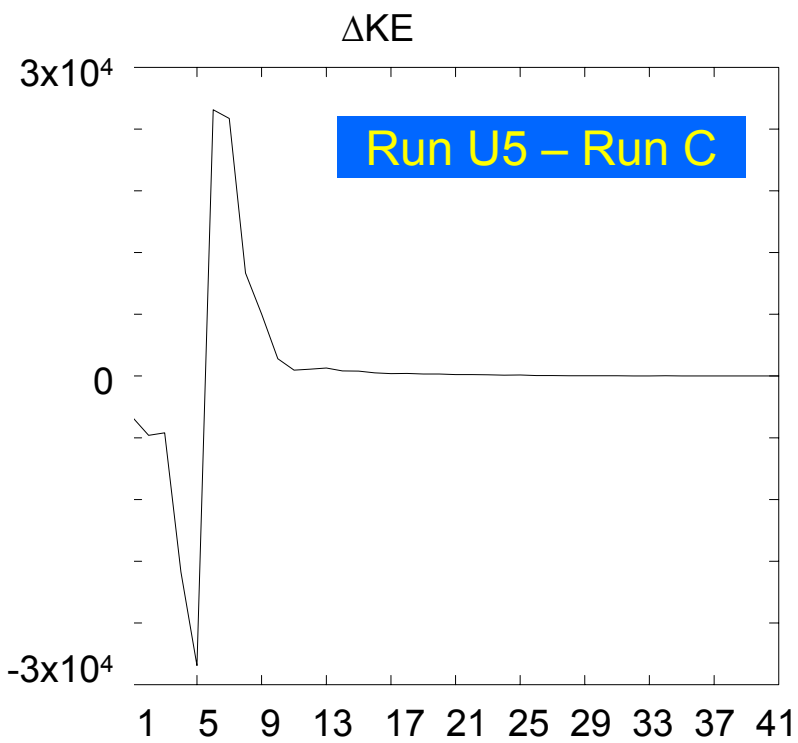
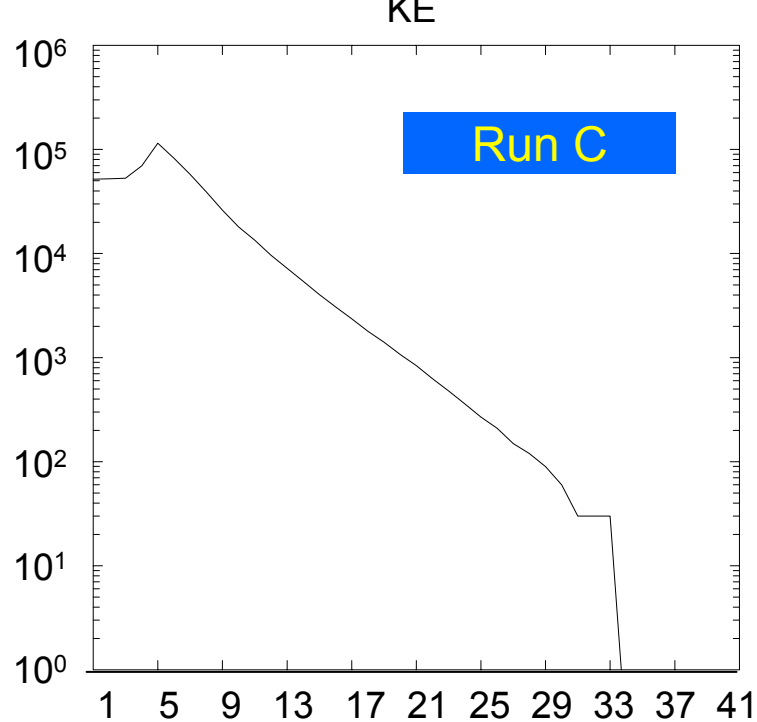
Contours: control run

Shading: anomaly

Vertically-integrated budget of zonal momentum



Spectrum of kinetic energy as a function of zonal wavenumber



Baroclinic lifecycle experiments

$\text{Log}_e(\text{kinetic energy})$ as a function of time (days) and wavenumber during an adiabatic baroclinic lifecycle running under background conditions of:

Run C

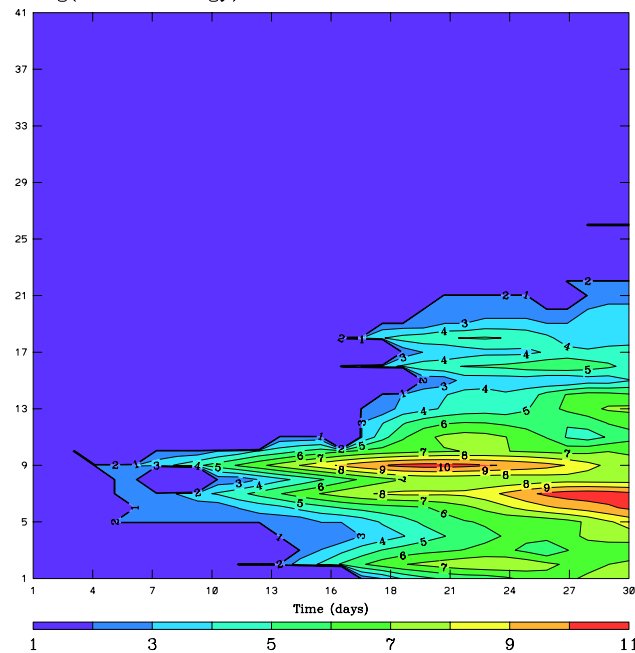
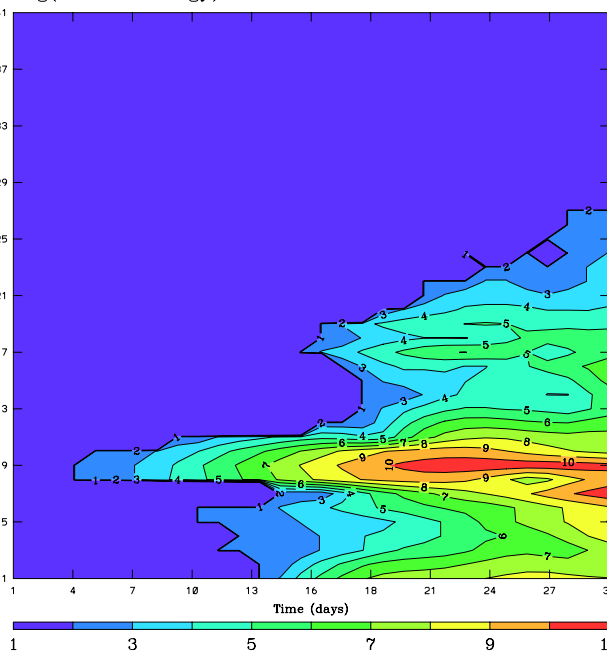
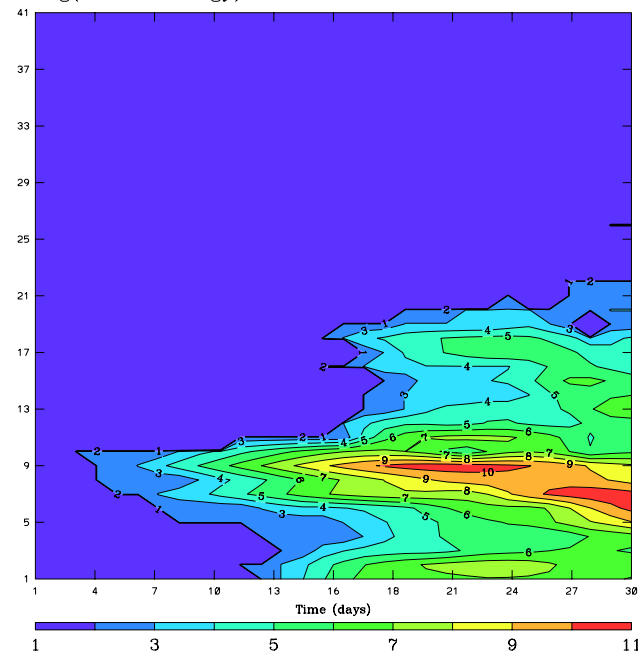
Run U5

Run E5

Log(kinetic energy) Run C

Log(kinetic energy) Run U5

Log(kinetic energy) Run E5



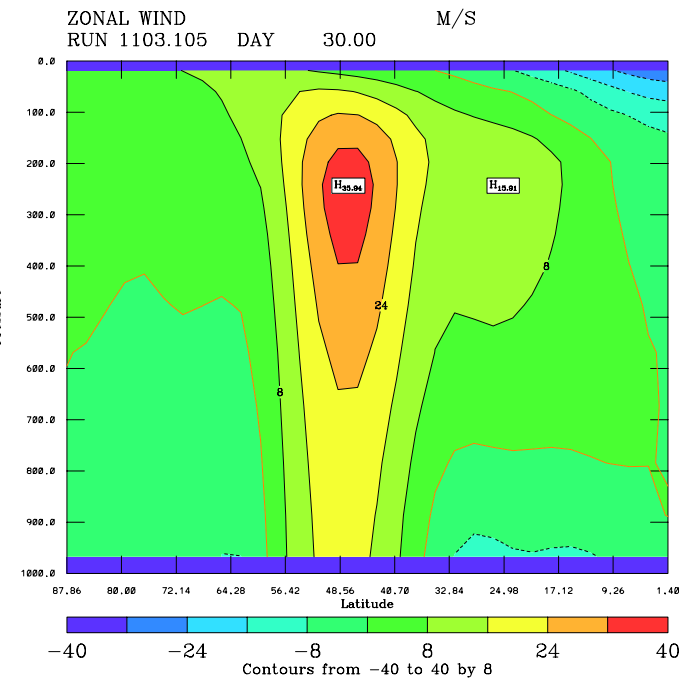
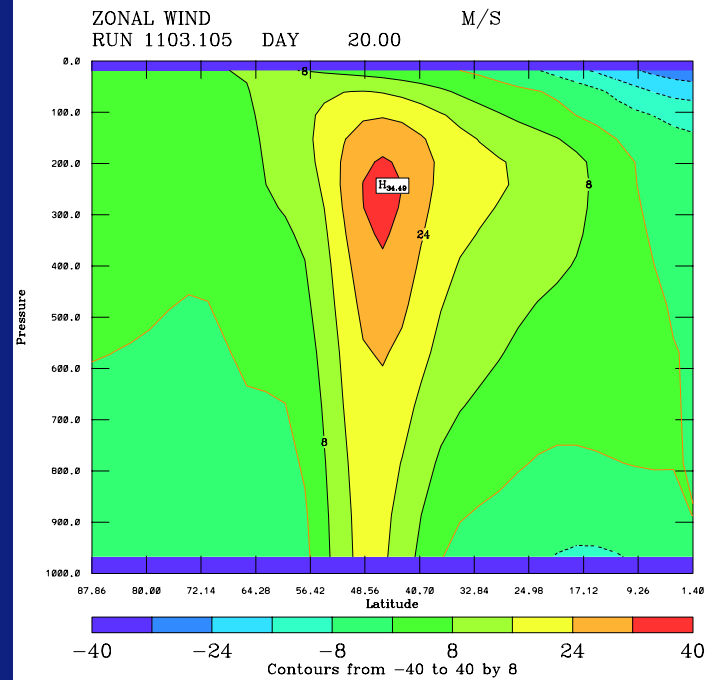
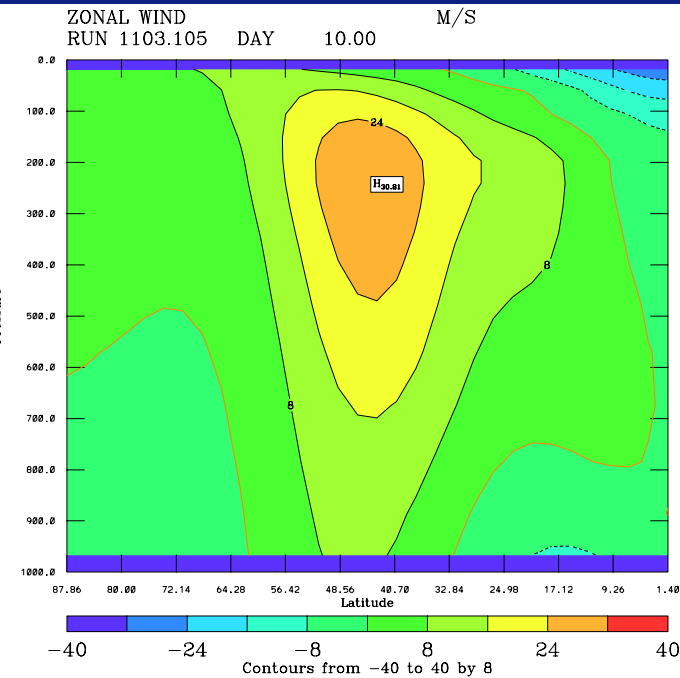
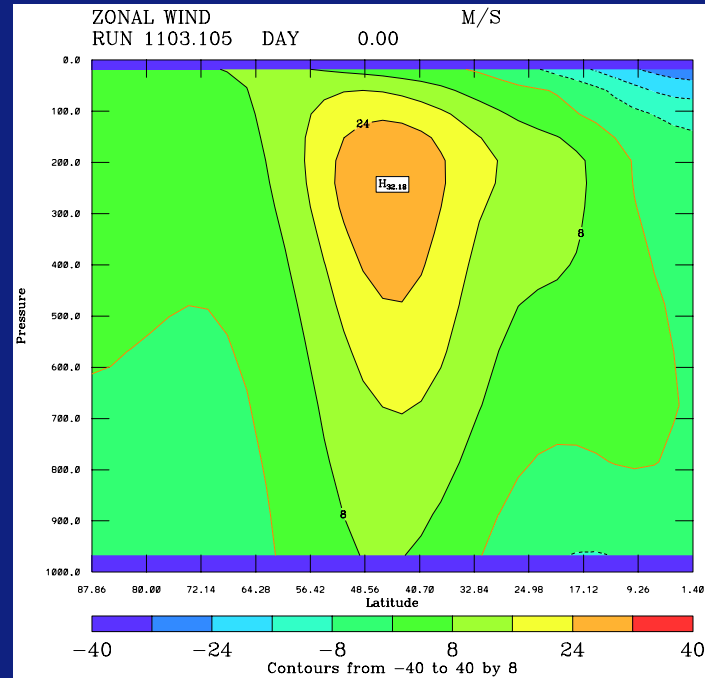
Zonal wind evolution during baroclinic lifecycle

Zonal mean zonal wind

(N. Hemisphere only)

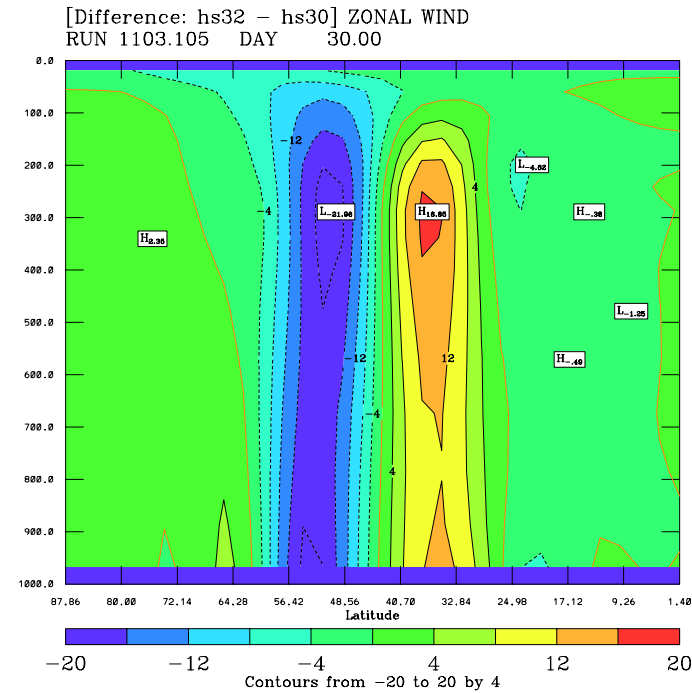
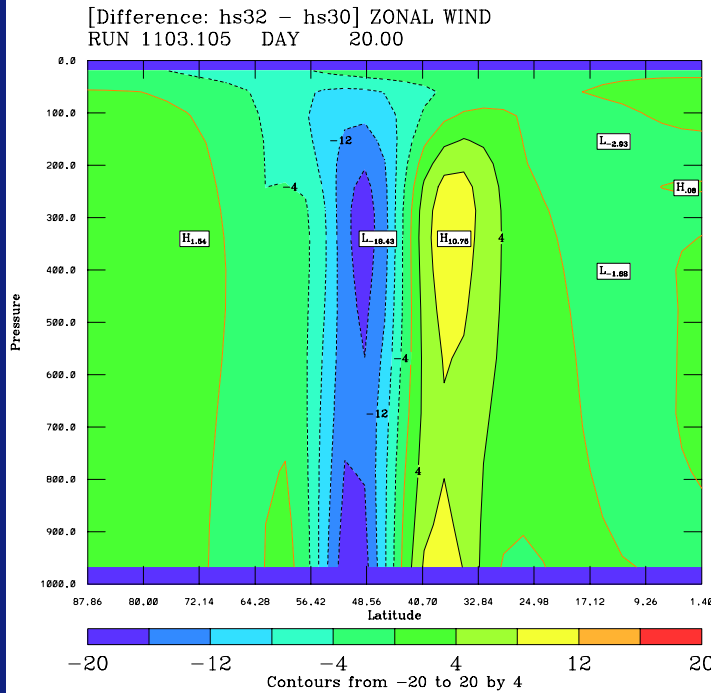
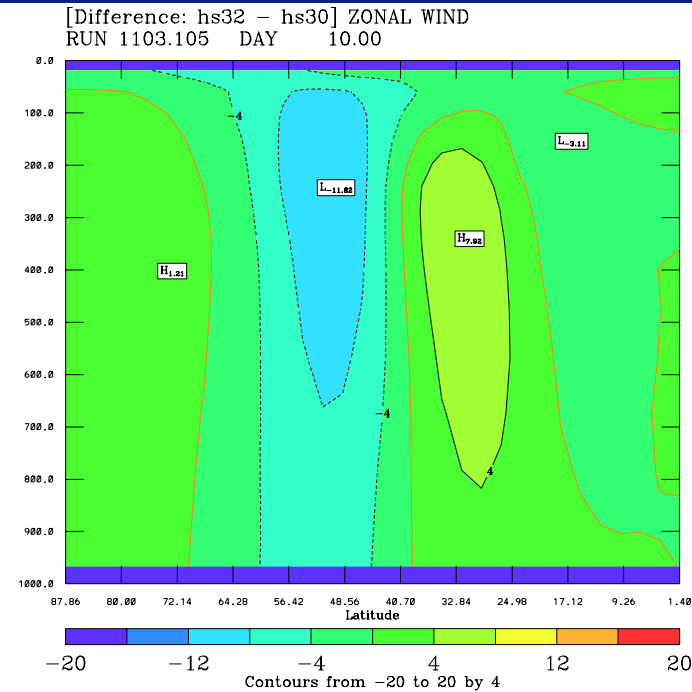
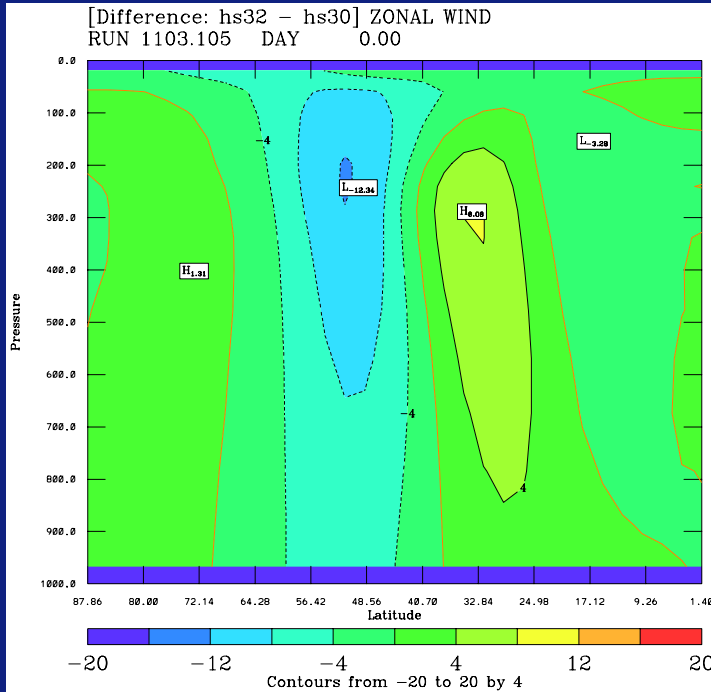
at days 0, 10, 20, 30

under the background conditions of Run C.



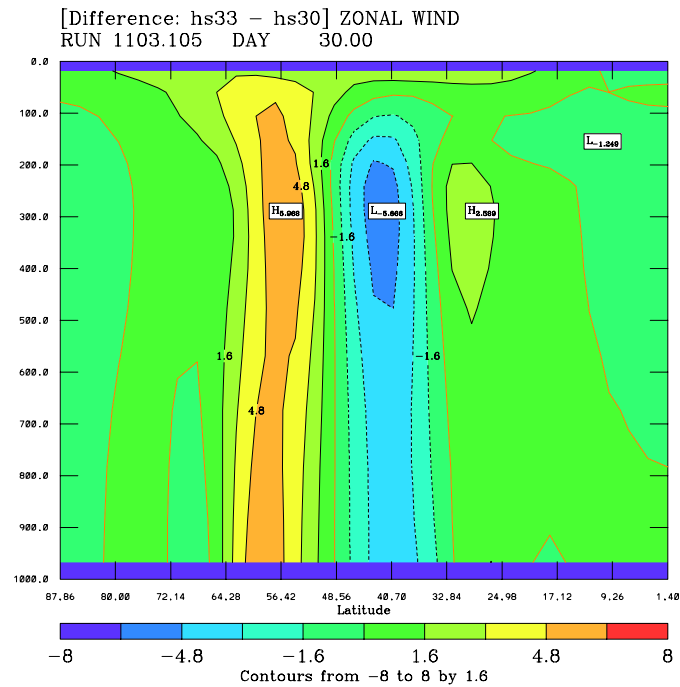
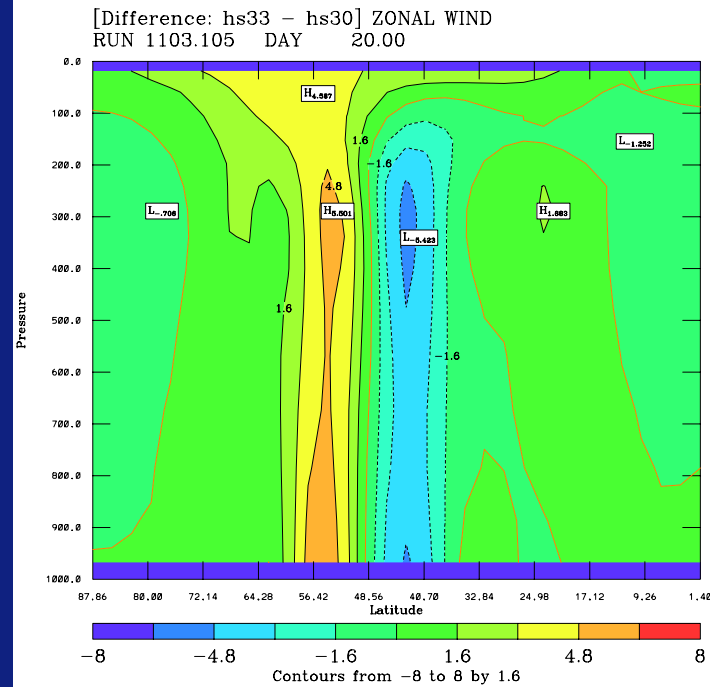
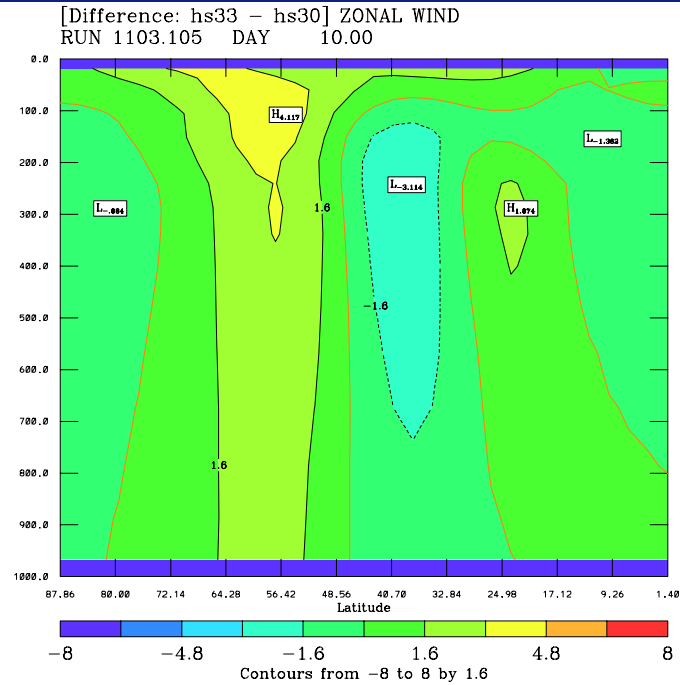
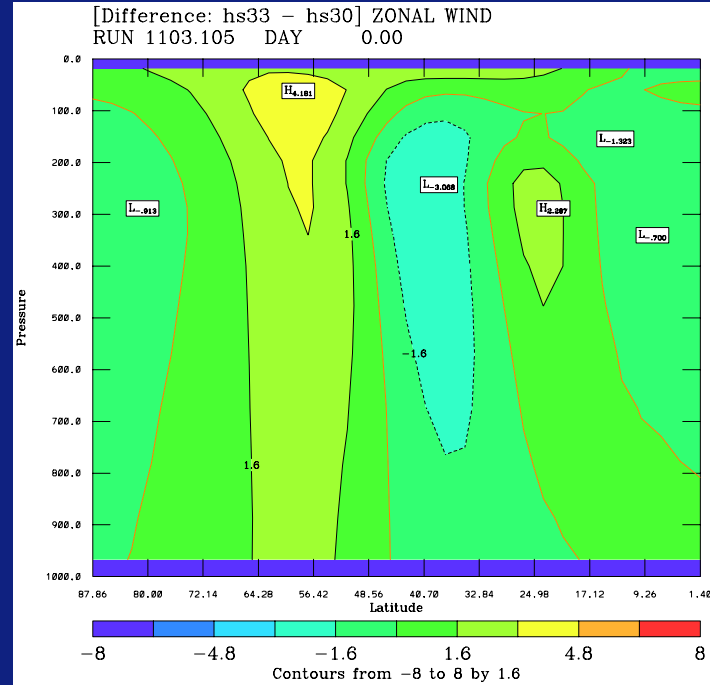
Zonal wind

Difference between baroclinic lifecycle runs under conditions of Run U5 and Run C



Zonal wind

Difference between baroclinic lifecycle runs under conditions of Run E5 and Run C



Conclusions

- Analysis of NCEP zonal winds reveals that when the sun is more active the sub-tropical jets are weaker and positioned nearer the poles.
- This signal is qualitatively similar to the results of GCM simulations with enhanced solar UV (and ozone).
- In a simplified GCM imposed stratospheric warming, and associated lowering of the tropopause, weakens the jets and storm-track eddies.
- Equatorial stratospheric warming displaces the jets polewards while uniform or polar warming displaces them markedly equatorwards.
- Baroclinic lifecycle runs show that baroclinic waves reinforce the zonal wind anomalies.