Potential of satellite spectral solar irradiance (SSI) measurements in ground-based remote sensing of atmospheric aerosols and trace gases

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Smoke from Pacific Northwest Fires
Retrieved from Suomi-NPP OMPS observations
Image from NASA Earth Observatory

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## 2017 Decadal Survey Recommended NASA Priorities

### Designated

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<tr>
<th>Categories</th>
<th>Targeted Observable</th>
<th>Science/Applications Summary</th>
<th>Candidate Measurement Approach</th>
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<td><strong>Designated</strong></td>
<td><strong>Aerosols</strong></td>
<td>Aerosol properties, aerosol vertical profiles, and cloud properties to understand their direct and indirect effects on climate and air quality</td>
<td>Backscatter lidar and multi-channel/multi-angle/polarization imaging radiometer flown together on the same platform</td>
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<td><strong>Earth System Explorer</strong></td>
<td><strong>Clouds, Convection, &amp; Precipitation</strong></td>
<td>Coupled cloud-precipitation state and dynamics for monitoring global hydrological cycle and understanding contributing processes</td>
<td>Radar(s), with multi-frequency passive microwave and sub-mm radiometer</td>
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<td><strong>Incubation</strong></td>
<td><strong>Mass Change</strong></td>
<td>Large-scale Earth dynamics measured by the changing mass distribution within and between the Earth’s atmosphere, oceans, ground water, and ice sheets</td>
<td>Spacecraft ranging measurement of gravity anomaly</td>
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<td><strong>Venture</strong></td>
<td><strong>Surface Biology &amp; Geology</strong></td>
<td>Earth surface geology and biology, ground/water temperature, snow reflectivity, active geologic processes, vegetation traits and algal biomass</td>
<td>Hyperspectral imagery in the visible and shortwave infrared, multi- or hyperspectral imagery in the thermal IR</td>
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<td><strong>Surface Deformation &amp; Change</strong></td>
<td>Earth surface dynamics from earthquakes and landslides to ice sheets and permafrost</td>
<td>Interferometric Synthetic Aperture Radar (InSAR) with ionospheric correction</td>
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Outline

- Aerosol retrieval from satellite observations
- Ground-based Aerosol remote sensing using sun photometers
- Calibration of sun photometers and relationship to SSI
- The need to use satellite SSI for ground instrument calibration
- Conclusions
Basic Concepts of Satellite Aerosol Retrieval Algorithms

\[ \rho_{\text{toa},\lambda} \sim \rho_{\text{surf},\lambda} + \rho_{\text{atm},\lambda} \]

Where, \( \rho_{\text{toa},\lambda} \) is the spectral reflectance measured at the top of the atmosphere (TOA)

\( \rho_{\text{surf},\lambda} \) is the component of the TOA reflectance contributed by the surface

\( \rho_{\text{atm},\lambda} \) is the component of the TOA reflectance from the atmosphere (or path radiance)

- **Over ocean**, \( \rho_{\text{surf},\lambda} \) is known (550-2130nm)
  - \( \rho_{\text{atm},\lambda} \sim \tau_{\lambda} \times P_{\lambda} (\theta) \); the spectral information can be used to derive information on aerosol type, which is then used to derive the aerosol loading.

- **Over land**, \( \rho_{\text{surf},\lambda} \) is complex
  - if \( \rho_{\text{atm},\lambda=2.1} \sim 0 \), then \( \rho_{\text{surf},\lambda=2.1} \) can be estimated
  - if \( \rho_{\text{surf},\lambda} \sim f (\rho_{\text{surf},\lambda=2.1}) \) then \( \rho_{\text{atm},\lambda} \) can be derived
  - \( \tau_{\lambda} \) can be derived assuming an aerosol model, i.e \( P_{\lambda} (\theta) \); if there are relationships between the spectral reflectances at different wavelengths.
Dubovik et al., 2002, JAS paper on aerosol optical properties (p. 591):

Modeling the aerosol effects on atmospheric radiation, by solving the radiative transfer equation, requires the following aerosol optical properties:

• aerosol optical thickness $\tau_\lambda$ (loading);

• phase function $P (\Theta ; \lambda)$ (angular dependence of light scattering),

• single-scattering albedo $\omega_{0\lambda}$ (ratio of scattering to scattering + absorption).
Figure 3. Summary of spectral optical properties of the (a–c) four smoke and (d–f) two dust aerosol optical models studied. Panels show spectral AOD (Figures 3a and 3c), spectral SSA (Figures 3b and 3e), and spectral ASY (Figures 3c and 3f). Paler to stronger colors are used to indicate models with weaker to stronger levels of absorption.
Models used for aerosol retrieval from satellite observations are often based on aerosol measurements from the ground, predominantly the Aerosol Robotic Network (AERONET) of ground-based sun photometers/sky radiometers. These models are often used to develop look-up tables (LUT) to facilitate aerosol retrieval.
Ground-based aerosol remote sensing based on direct spectral solar irradiance: AERONET example

https://aeronet.gsfc.nasa.gov
Ground-based AERONET used to validate satellite aerosol retrieval

MAPSS (Multi-sensor Aerosol Products Sampling System) http://giovanni.gsfc.nasa.gov/mapss/

AERONET, MODIS, MISR, OMI, POLDER, CALIOP, SeaWiFS, VIIRS

Petrenko et al., 2012, AMT; Petrenko and Ichoku, 2013, ACP
Ship-based MAN used to validate satellite aerosol retrieval

Frequency distributions of aerosol optical depth (AOD) differences between various satellite sensors and sun-photometer.

Smirnov et al., 2017, SPIE
Aerosol remote sensing based on Beer-Bouguer-Lambert’s law. Simply:

\[ I(\lambda) = I_0(\lambda) \exp[-\tau(\lambda)/\cos \theta] \]

Where,
- \( I_0(\lambda) \) is the extraterrestrial flux at wavelength \( \lambda \),
- \( I(\lambda) \) is the flux reaching the ground, \( \theta \) is the solar zenith angle, and
- \( \tau(\lambda) \) is the total optical depth

**Dutton et al., 1994, JGR states:**

“The need for absolute irradiance measurements is eliminated for \( \tau(\lambda) \) computation because the linear response of the sunphotometer is such that a linear conversion constant cancels from both sides of the equation, and the instrument output voltages, \( V(\lambda) \) and \( V_0(\lambda) \) replace \( I(\lambda) \) and \( I_0(\lambda) \).”

**This means:** \( V(\lambda) = k \times I(\lambda) \) and \( V_0(\lambda) = k \times I_0(\lambda) \)

Where \( k \) is a convolution of various instrument specific parameters (filter transmission, field of view, detector sensitivity, etc.)
Thus, for direct sun measurement using sun photometers:

\[ V_\lambda = V_{0\lambda} D^{-2} \exp(-\tau_\lambda M) \]

\[ V_w = V_{0w} D^{-2} \exp[-\tau_w M - k (WM)^b] \]

where, for each channel (wavelength \( \lambda \) in microns),
\( V_\lambda \) = the signal measured by the instrument at wavelength \( \lambda \),
\( V_{0\lambda} \) = the extraterrestrial signal at wavelength \( \lambda \),
\( D \) = Earth-Sun Distance in Astronomical units at time of observation,
\( \tau_\lambda \) = total optical thickness (\( \tau_\lambda = \tau_{a\lambda} + \tau_{R\lambda} + \tau_{O3\lambda} \)) at wavelength \( \lambda \),
\( \tau_{a\lambda} \) = aerosol optical thickness (AOT) at wavelength \( \lambda \),
\( \tau_{R\lambda} \) = Rayleigh (air) optical thickness at wavelength \( \lambda \),
\( \tau_{O3\lambda} \) = Ozone optical thickness at wavelength \( \lambda \),
\( M \) = the optical air mass
\( W \) = vertical water vapor column thickness
\( k \) and \( b \) are instrument constants numerically derived for the 936 nm filter.

Ichoku et al., 2002, JGR, Microtops
Sun-photometer calibration approaches:

Langley calibration under pristine conditions (Mauna Loa)
\[
\ln(V_\lambda) = \ln(V_{0\lambda} D^{-2}) - \tau_\lambda M
\]

Or

Transfer calibration from a Langley-calibrated Instrument
\[
\frac{V_\lambda}{V'_{0\lambda}} = \frac{V_{0\lambda}}{V'_{0\lambda}}
\]

Ichoku et al., 2002, JGR, Microtops
Transfer Calibration of five Microtops sun-photometers against AERONET at NASA/GSFC

**Experiences**

- Proper calibration takes a long time
- Pre- and Post-deployment calibration always required
- Uncertainty persists (why do the $V_0$s vary over time?)

Ichoku et al., 2002, JGR, Microtops
AOT values can differ significantly depending on calibration method

Ichoku et al., 2002, JGR, Microtops
TOA-measured SSI seems more realistic than modeled

Harder et al., 2005, Solar Physics
Herman et al., 2015, AMT paper on Pandora (p 3409):
The laboratory-calibrated Pandora TCO retrieval algorithm uses an external solar reference spectrum derived from a combination of the Kurucz spectrum (wavelength resolution \(\lambda/\Delta\lambda = 500\ 000\)) radiometrically normalized to the lower-resolution shuttle Atlas-3 SUSIM spectrum (Van Hoosier, 1996; Bernhard et al., 2004).
Herman et al., 2015, AMT paper on Pandora (p 3409): The use of a well-calibrated top-of-the-atmosphere spectrum convolved with the laboratory-measured spectrometer slit function derived for each pixel permits derivation of ozone amounts without resorting to either a Langley calibration approach or calibration transfer from a standard instrument. The core slit function is known to within 1%, which propagates into an ozone error of less than 1%.

Smirnov, 2018, personal communication: This statement is correct, but only if we are sure that SSI is known with an uncertainty of less than 0.5%.
Coddington et al., 2014, BAMS paper on Solar Irradiance:

Table 1. Measurement requirements established for the TSIS TIM and SIM instruments that are driven by the need to understand Earth’s climate response to solar variability, for separating natural from anthropogenic climate forcing effects, and for the monitoring and interpretation of the variability in wavelength-dependent processes induced by changes in Earth’s surface and atmosphere.

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<tr>
<th>Parameter</th>
<th>TSI CDR requirement</th>
<th>SSI CDR requirement</th>
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<tr>
<td>Absolute accuracy</td>
<td>0.01%</td>
<td>0.2%</td>
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<tr>
<td>Stability</td>
<td>0.001% yr⁻¹</td>
<td>0.05% yr⁻¹ (λ &lt; 400 nm)</td>
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<tr>
<td>Relative precision</td>
<td>0.001%</td>
<td>0.01% (λ &gt; 400 nm)</td>
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Take Home Message: If SSI measurements from TSIS and future missions meet these requirements, they should be seriously considered for calibration of Sun-pointing ground-based instruments.
Acknowledgements

The TSIS-1 Project
The AERONET Project
The Pandora Global Network (PGN) Project
References


