

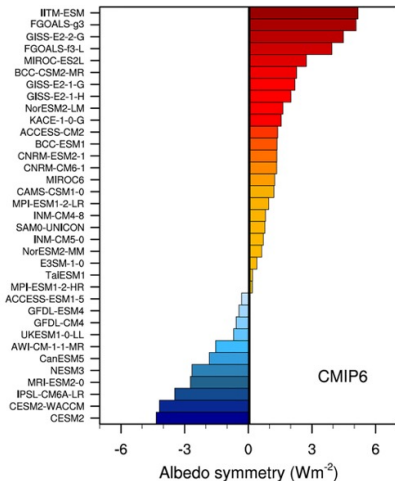
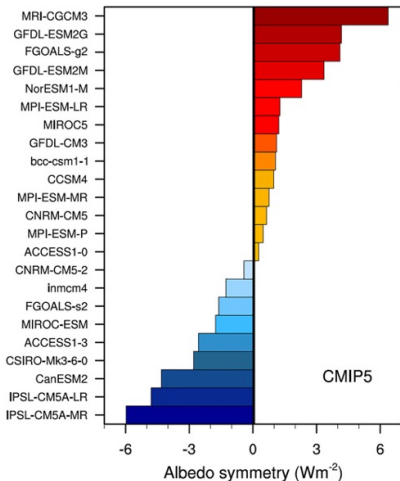


# Evidence for Hemispheric Spectral Albedo Inequality

William D. Collins<sup>1,2</sup> and Daniel R. Feldman<sup>1</sup>

*<sup>1</sup>Berkeley Lab and <sup>2</sup>University of California, Berkeley*

# Earth's Reflected Shortwave Symmetry – not Seen in ESMs

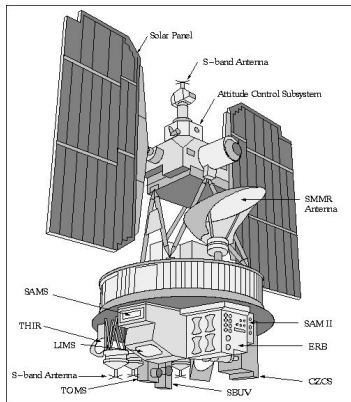


Nimbus-7 was last satellite with working near-IR broadband radiometers

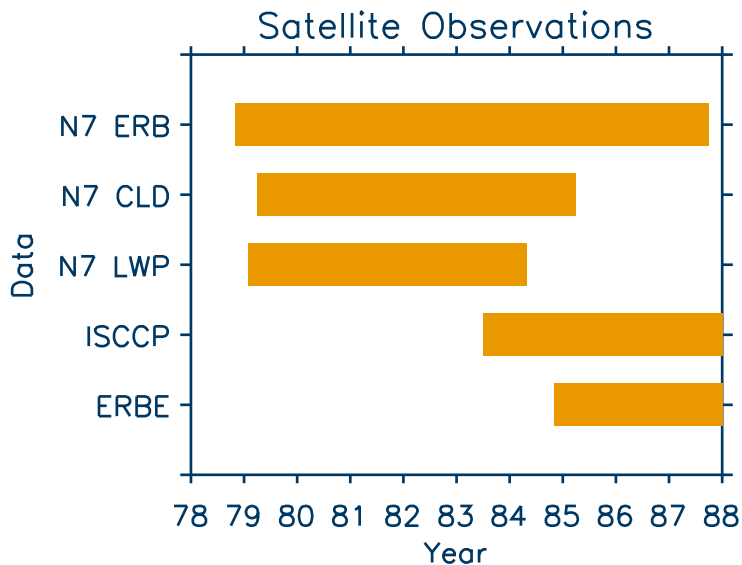


# Nimbus-7 Instrument Suite

- ▶ ERBS (Earth Radiation Budget Sensor)
- ▶ CZCS (Coastal-Zone Color Scanner)
- ▶ LIMS (Limb Infrared Monitoring of the Stratosphere)
- ▶ SAM II (Stratospheric Aerosol Measurement II)
- ▶ SAMS (Stratospheric and Mesospheric Sounder)
- ▶ SBUV (Solar Backscatter UV)
- ▶ SMMR (Scanning Multichannel Microwave Radiometer)
- ▶ THIR (Temperature-Humidity Infrared Radiometer)
- ▶ TOMS (Total Ozone Mapping Spectrometer)



# Timeline of Data for Earth Radiation Budget studies



## Relationship of Spectral to Broadband Albedos

The TOA broadband albedo in each hemisphere  $i = NH, SH$  is:

$$R_i = f_v R_{v,i} + (1 - f_v) R_{n,i}$$

This can be written as:

$$\begin{aligned} R_i &= \frac{\sum R_{b,i}}{2} \left[ 1 + \frac{\Delta R_{b,i}}{\sum R_{b,i}} (2 f_v - 1) \right] \\ &= \frac{\sum R_{b,i}}{2} [s_i] \\ &= R_i \left[ \frac{s_i}{s_i} \right] \end{aligned}$$

where

$$\begin{aligned} \sum R_{b,i} &= R_{v,i} + R_{n,i} \\ \Delta R_{b,i} &= R_{v,i} - R_{n,i} \\ s_i &= \left[ 1 + \frac{\Delta R_{b,i}}{\sum R_{b,i}} (2 f_v - 1) \right] \end{aligned}$$

# Implications of Spectral Interhemispheric Differences

These results show that

$$R_{NH} \simeq R_{SH}$$

can be true even if

$$s_{NH} \neq s_{SH}$$

Let us write

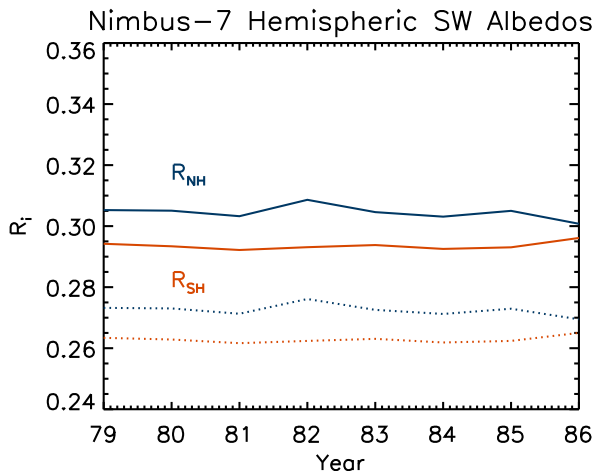
$$s_i = [1 + r_i (2 f_v - 1)]$$
$$r_i = \frac{\Delta R_{b,i}}{\Sigma R_{b,i}}$$

Note that

$$-1 \leq r_i \leq 1 \quad r_i = \begin{cases} -1 \Rightarrow R_{v,i} = 0 \Rightarrow \text{max Surf. heating} \\ 1 \Rightarrow R_{n,i} = 0 \Rightarrow \text{max Atm. heating} \end{cases}$$

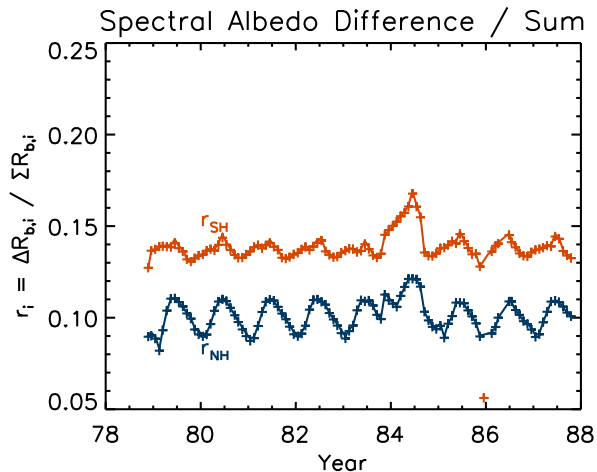
**What does Nimbus-7 show?**

# Nimbus-7 Hemispheric Shortwave Albedos



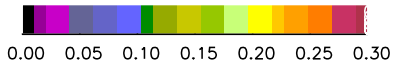
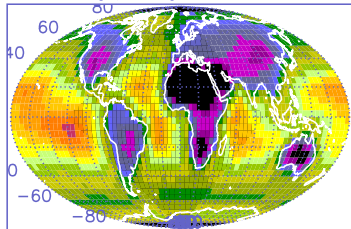


# Spectral Albedo Difference / Sum

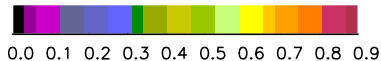
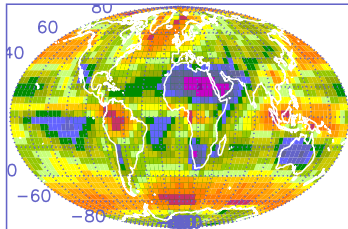


# Time-Mean Spectral Albedo Difference / Sum & Cloud Amount

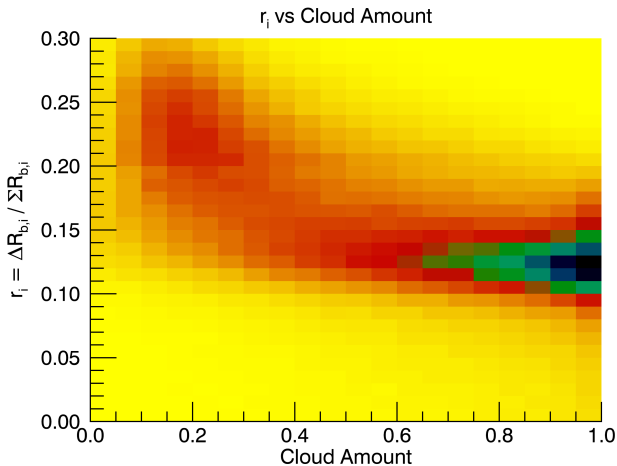
Mean  $\Delta R_b / \Sigma R_b$  for 1980–1984



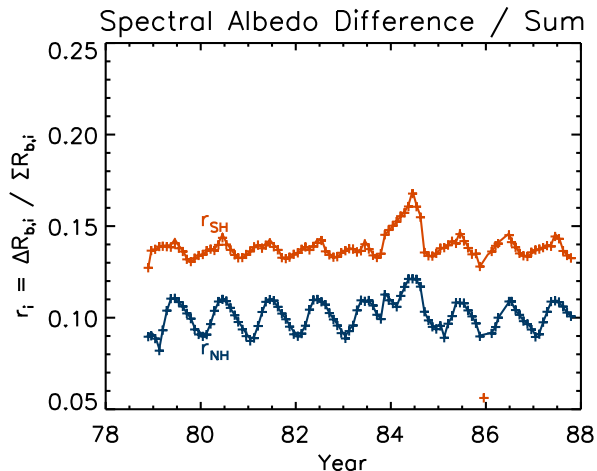
Mean Cloud Amount for 1980–1984



# Daily Spectral Albedo Difference / Sum & Cloud Amount over Oceans

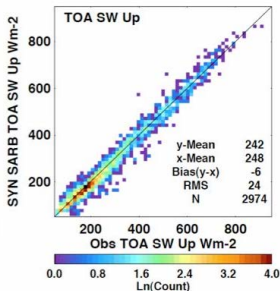
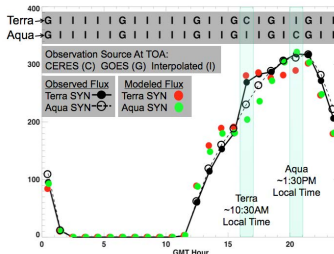
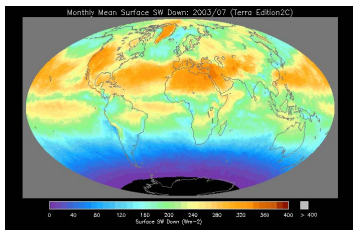


# Can the Nimbus-7 Results be Tested with Recent Data?



# CERES SYN Product (2003–2021)

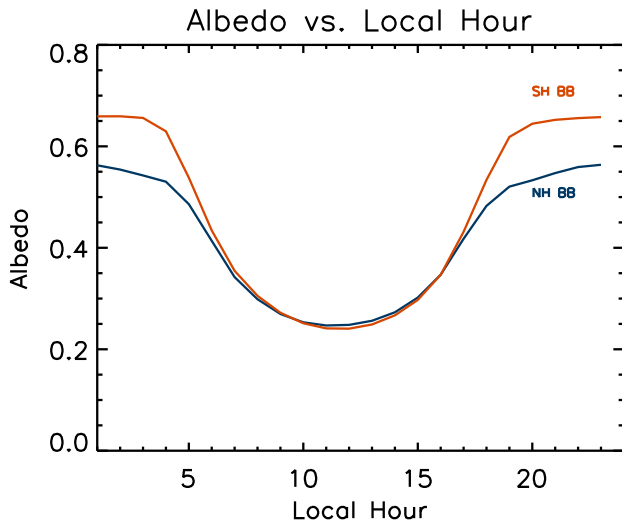
Global Hourly Gridded TOA BB, BB clear, and Spectral Fluxes:



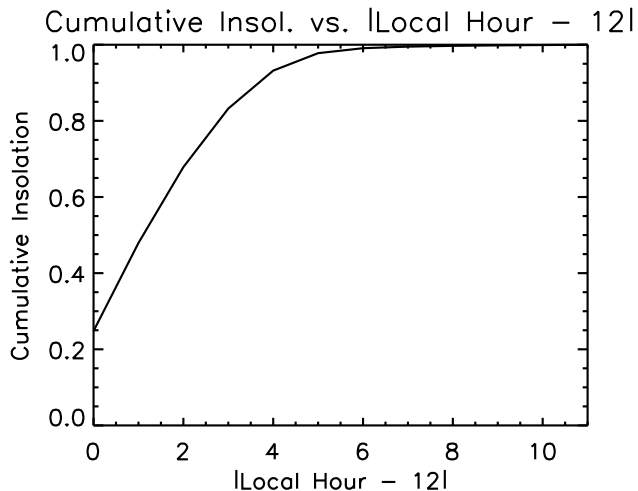
# Hemispheric Albedos from CERES Products

Hemisphere	EBAF	SYN
NH	0.291	0.285
SH	0.291	0.285

# Variation of Hemispheric Albedo with Local Hour

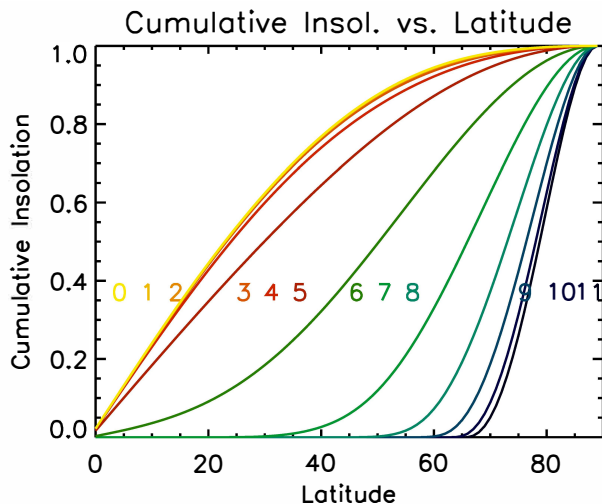


## Cumulative Insolation vs. Local Hr Range about Noon





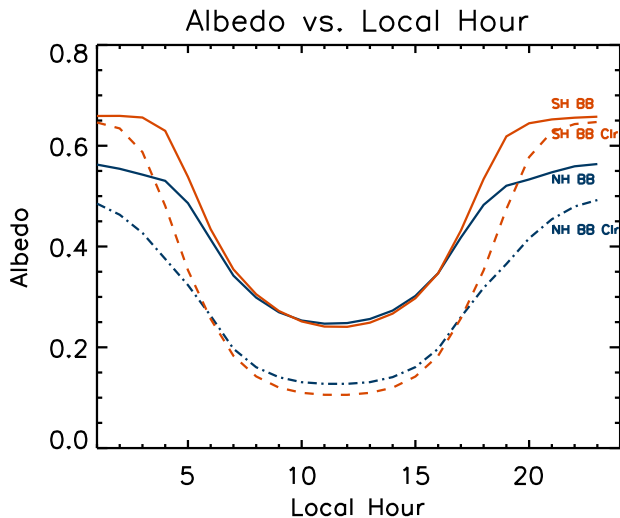
# Cumulative Insolation by Local Hr vs Latitude



# Hemispheric Albedos from CERES Products

Hemisphere	EBAF	SYN	Noon SYN	Nimbus-7
NH	0.291	0.285	0.247	0.304
SH	0.291	0.285	0.241	0.294

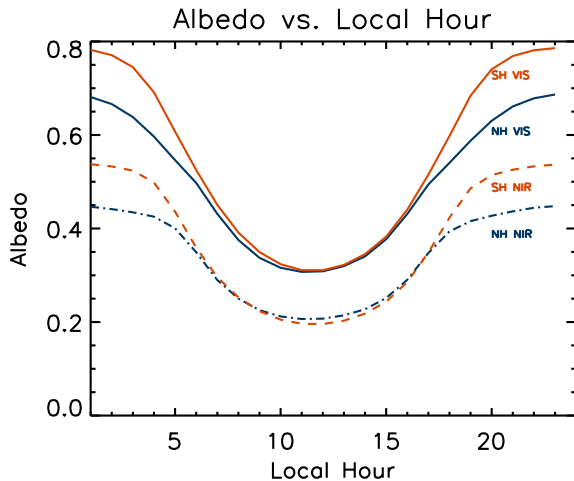
# Variation of Clear-sky Hemispheric Albedo with Local Hour



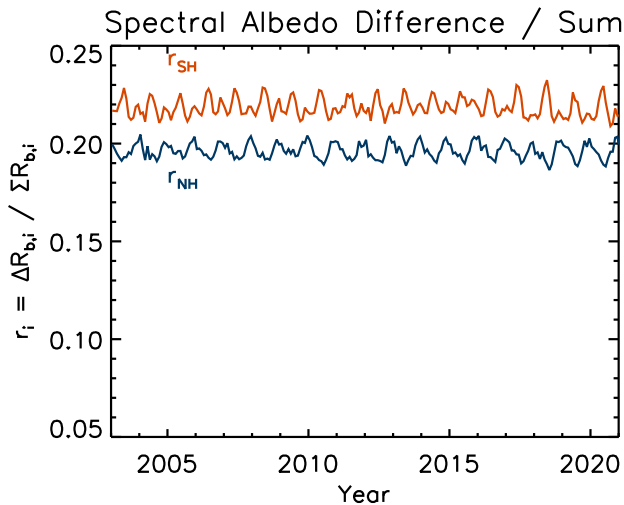
# Hemispheric Albedos from CERES Products

Hemisphere	EBAF	SYN	Noon SYN	Spectral SYN
NH	0.291	0.285	0.247	0.292
SH	0.291	0.285	0.241	0.293

# Variation of Spectral Hemispheric Albedo with Local Hour

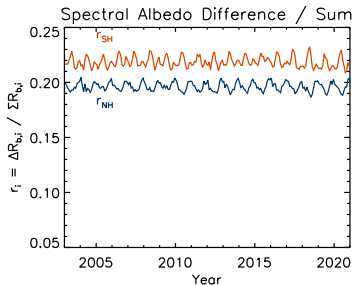


# CERES SYN Spectral Albedo Difference / Sum

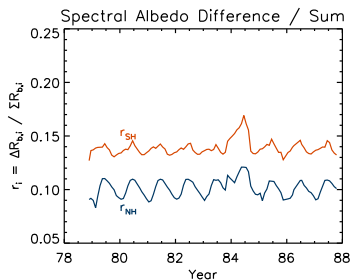


# Comparison to Nimbus-7 Spectral Albedo Difference / Sum

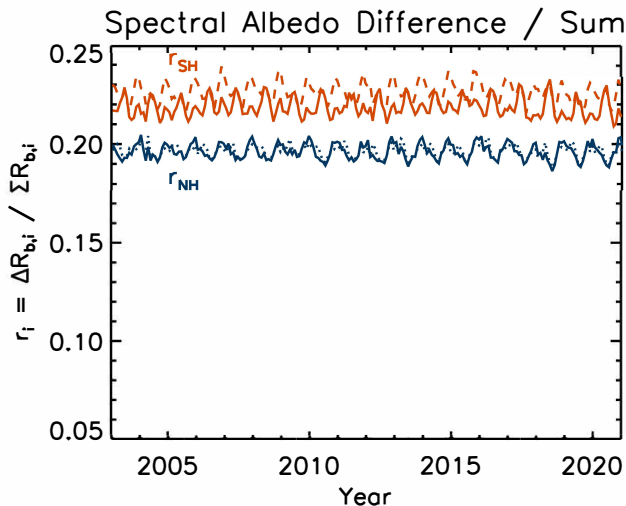
CERES SYN:



NIMBUS-7:



# CERES SYN Spectral Albedo Diff. / Sum @ Local Noon





## Open Issues

- ▶ What are the implications of these findings for possible mechanisms (–ve feedbacks) to maintain

$$R_{NH} \simeq R_{SH}$$

- ▶ Since  $f_v = 0.497$ , the spectral multipliers on  $R_i$  are insensitive to the spectral partitioning between visible and near-IR:

$$\frac{\partial s_i}{\partial r_i} = 2 f_v - 1 \simeq 0$$

What are the implications (if any)?

- ▶ The spectral factors are related to the albedo ratio  $\mathcal{R}_i = R_{n,i}/R_{v,i}$  (Collins, 1998) by:

$$r_i = \frac{1 - \mathcal{R}_i}{1 + \mathcal{R}_i}$$

This means  $r_i$  will differ between obs and models for cloudy conditions over oceans. What are the implications?